

1 **MODELING AND MEASURING FERRY BOAT WAKE**
2 **PROPAGATION**

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4 **ABSTRACT**

5 Surface waves generated by vessels were studied for two classes of ferries in the
6 Washington State Department of Transportation (WSDOT) system. The vessel
7 wakes were measured using a suite of free-drifting GPS buoys. The wakes were
8 modeled using an analytical potential flow model, which has a strong dependence
9 on vessel Froude number and vessel aspect ratio. Both the measurements and
10 the model are phase-resolving, and thus direct comparisons are made, in addition
11 to comparing bulk parameters. The model has high fidelity for the parameters
12 describing the operating conditions for WSDOT vessels, successfully reproducing
13 the measured difference between the two vessel classes and the dependence on
14 vessel speed for each vessel class.

15 **Keywords:** Free Surface, Ship Wake, Kelvin Wake.

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INTRODUCTION

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Washington State Ferries (WSF) operates a fleet of 22 vessels in the Puget Sound region of Washington State, USA, and British Columbia, Canada. It is the largest ferry system in the United States and services the highest volume of automobiles of any ferry service in the world. The Seattle-Bremerton route, servicing 2.33 million passengers and 642,000 vehicles per year (Smith 2013), is the primary subject of the current study. In a section of this route named Rich Passage, denoted by the box in Fig. ??, the channel narrows to 600m (?). Due to the narrow channel and steep bathymetry of the region, vessel wake-induced erosion is a concern for the surrounding shoreline.

A vessel operating protocol (power limit) of 8.2 m/s (16 knots) is presently in place to reduce wake-induced morphological effects within the channel (Klann 2002). The WSF vessels that serve the Seattle-Bremerton route are known as Issaquah Class and Super Class vessels. While of qualitatively similar size, the Super Class vessels are longer, narrower, and heavier than the Issaquah Class vessels. The protocol is applied to the Issaquah Class only, because the Issaquah Class wakes have been classified, visually, as much larger than the Super Class wakes. This paper seeks to quantify the differences between the wakes of the two vessel classes through the use of field data and an analytical model.

Rich Passage has been the subject of a comprehensive wake and shoreline study conducted upon the introduction of the Rich Passage I, a passenger-only fast ferry that served the Seattle-Bremerton route for Kitsap Transit. The study, conducted by Golder and Associates, concluded that the Rich Passage I meets wake height guidelines specified by

$$H = \begin{cases} H < 0.2 & T \leq 3.5 \\ H < 1.16T^{-1.4} & T > 3.5 \end{cases} \quad (1)$$

41 where H is the the wake height in meters measured 300m from the sailing line of
42 the vessel, and T is the wake period in seconds. This criterion is based on the
43 observation that low-frequency waves contain more erosive potential because of
44 their higher energy flux, and therefore have a more restrictive associated height
45 limit. The morphological shoreline response was largely inconclusive (Cote et al.
46 2013).

47 WSF commissioned a similar study of the car ferry fleet in 1985 to measure
48 vessel wakes (Nece et al. 1985). This study involved Issaquah, Superclass, and
49 Evergreen class models in open water in Puget Sound. No modeling efforts ac-
50 companied this study, and no curve fitting or statistical analysis were performed.
51 The results establish qualitatively that the Super class vessel wakes grow at a
52 much slower rate as vessel speed increases when compared to the Issaquah class.
53 However, noise in wake data and difficulties in establishing the buoy-boat dis-
54 tance make the conclusions difficult to quantify. Despite these shortcomings, it
55 was found that the data from this study could be used to supplement the field
56 data from the current study. Uncertainty in wake height due to natural varia-
57 tion in time-dependent surface wave conditions dominated uncertainty due to the
58 buoy-boat distances.

59 The more general objective of predicting boat wakes has generally been ap-
60 proached as either a computational fluids problem, or as an empirical quantifi-
61 cation of wave height. Previous modeling efforts of phase resolved vessel wakes
62 have been confined to domains of only a few boat lengths due to computational
63 requirements (Raven 1998), which provide limited information about potential
64 shoreline impact. Sophisticated near-vessel models are primarily used to model
65 wave breaking and wave making resistance for naval architectural applications
66 (Landrini et al. 1999; Huang and Yang 2016). Phase-averaged models have been
67 successfully implemented (Cote et al. 2013), but the resultant wake structure

68 is lost. Other phase-resolved models have focused on vessels with transcritical
69 Froude numbers in shallow water (Kofoed-Hansen et al. 1999). Empirical models
70 have been used to great success within the bounds of the space sampled (Sorensen
71 1997; Kriebel et al. 2003; Ng and Bires)

72 The analytic study of ship wakes was first considered mathematically in 1887,
73 when Lord Kelvin showed that the disturbance caused by a point pressure source
74 moving across a fluid surface forms a predictable structure which now bears his
75 name (Thomson 1887). Subsequent work focused on the geometry and structure of
76 wakes, with extensions for arbitrary motion of the pressure point (Stoker 1957).
77 Significant gains have been made more recently, with results detailing surface
78 velocity fields (Yun-gang and Ming-de 2006), the inclusion of viscous forces (Lu
79 and Chwang 2007), and finite vessel dimensions (Darmon et al. 2014; Benzaquen
80 et al. 2014).

81 The impact of boat wakes on shoreline, and mitigation thereof, has been a
82 continuous concern for caretakers of waterways, as well as a subject of many
83 previous studies (Glamore 2008; Aage et al. 2003; Verhey and Bogaerts 1989). In
84 particular, work has been done on the sediment transport associated with vessel-
85 generated waves (Velegrakis et al. 2007).

86 The effects of bow wakes, which are known to have complex hull-dependent
87 forms (Noblesse et al. 2013) are not considered in this study, though results suggest
88 that they may dominate the vessel-generate wave field for low Froude number.

89 **METHODS**

90 **Analytical Model**

91 The first analytic study of wake patterns was published in 1887, when Lord
92 Kelvin showed that the disturbance caused by a point pressure source moving
93 across a fluid surface forms a predictable structure which now bears his name
94 (Thomson 1887). Subsequent work focused on the geometry and structure of

95 wakes, with extensions for arbitrary motion of the pressure point (Stoker 1957).
 96 A closed form expression for the sea surface height caused by a finite pressure
 97 distribution moving across a fluid surface was recently developed (Benzaquen et al.
 98 2014; Darmon et al. 2014). The equations developed by Darmon, Benzaquen, et
 99 al. are the basis for the analytical modeling presented below.

100 The analytical model must be confined to potential flow theory in order to
 101 yield tenable solutions for the sea surface height. Therefore, the restrictions

$$102 \quad \mu = 0 \quad \text{and} \quad \vec{\omega} = \nabla \times \vec{u} = 0, \quad (2)$$

103 are required. Here, μ is the viscosity of the flow, \vec{u} is the flow field in Euclidean
 104 coordinate space, and $\vec{\omega}$ is the vorticity.

105 The sea surface height generated by a moving pressure distribution $p(x, y)$
 106 (Raphaël and deGennes 1996) is given by

$$107 \quad \zeta(x, y) = -\lim_{\epsilon \rightarrow 0} \int_0^\infty \int_0^\infty \frac{dk_x dk_y}{4\pi^2 \rho} \frac{k \hat{p}(k_x, k_y) e^{-i(k_x x + k_y y)}}{\omega(k)^2 - U^2 k_x^2 + 2i\epsilon U k_x} \quad (3)$$

108 where $\hat{p}(k_x, k_y)$ is the Fourier transform of the pressure distribution $p(x, y)$, ρ is the
 109 constant fluid density, $\omega(k)$ is the surface wave dispersion relation, $k = \sqrt{k_x^2 + k_y^2}$,
 110 and U is the pressure distribution's constant forward velocity. The parameter ϵ is a
 111 variable with units s^{-1} that approaches zero, ensuring that the radiation boundary
 112 condition is satisfied. For a derivation of Eq. (3), the reader is referred to Raphaël
 113 and deGennes (1996). Define a coordinate system such that the vessel is located
 114 at the origin, and is moving in the $-x$ -direction, with the sea surface displacement
 115 in the z -direction. The y -axis defines the neutral sea surface perpendicular to the
 116 sailing line of the vessel. Following the method developed by Darmon et al., the

117 nondimensionalization conditions for a boat of length b are:

$$\begin{aligned}
X &= \frac{x}{b} & Y &= \frac{y}{b} & K_X &= k_x b \\
K_Y &= k_y b & Z &= \frac{4\pi^2 \zeta}{b} & \hat{P} &= \frac{\hat{p}}{\rho g b^3} & \tilde{\epsilon} &= \epsilon \sqrt{\frac{b}{g}},
\end{aligned} \tag{4}$$

119 Nondimensionalization yields

$$\begin{aligned}
\frac{Z(X, Y)b}{4\pi^2} &= -\lim_{\tilde{\epsilon} \rightarrow 0} \int_{-\pi/2}^{\pi/2} \int_0^\infty \frac{dK_X dK_Y}{4\pi^2 b^2 \rho} \frac{K}{b} \rho g b^3 \hat{P}(K_x, K_y) e^{-i(K_X X + K_Y Y)} \\
&\quad \frac{gK}{b} - U^2 \frac{K_X^2}{b^2} + 2i\tilde{\epsilon} \sqrt{\frac{g}{b}} U \frac{K_X}{b} \\
Z(X, Y) &= -\lim_{\tilde{\epsilon} \rightarrow 0} \int_{-\pi/2}^{\pi/2} \int_0^\infty \frac{dK_X dK_Y}{b^4 \rho} \frac{K \rho g b^3 \hat{P}(K_x, K_y) e^{-i(K_X X + K_Y Y)}}{\frac{gK}{b} - U^2 \frac{K_X^2}{b^2} + 2i\tilde{\epsilon} \sqrt{\frac{g}{b^3}} U K_X} \\
&= -\lim_{\tilde{\epsilon} \rightarrow 0} \int_{-\pi/2}^{\pi/2} \int_0^\infty \frac{dK_X dK_Y K \hat{P}(K_x, K_y) e^{-i(K_X X + K_Y Y)}}{K - Fr^2 K_X^2 + 2i\tilde{\epsilon} Fr K_X}
\end{aligned} \tag{5}$$

121 An elliptic transform

$$122 \quad X = \check{R} \cos \check{\varphi} \quad Y = W \check{R} \sin \check{\varphi} \quad K_X = \check{K} \cos \check{\theta} \quad K_Y = W^{-1} \check{K} \sin \check{\theta}, \tag{6}$$

123 permits solutions for an elliptic Gaussian pressure distribution. Define an aspect

124 ratio $W > 0$ to be the ratio of the y to x characteristic dimensions of $P(x, y)$.

125 Applying this transformation,

$$\begin{aligned}
Z(\check{R}, \check{\varphi}) &= \\
&= -\lim_{\tilde{\epsilon} \rightarrow 0} \int_{-\pi/2}^{\pi/2} \int_0^\infty \frac{\check{K}}{W} d\check{K} d\check{\theta} \frac{\check{K} \sqrt{\cos^2 \check{\theta} + W^{-2} \sin^2 \check{\theta}} \check{P}(\check{K}) W e^{-i\check{K} \check{R} (\cos \check{\varphi} \cos \check{\theta} + \sin \check{\varphi} \sin \check{\theta})}}{\sqrt{\cos^2 \check{\theta} + W^{-2} \sin^2 \check{\theta}} - Fr^2 \check{K}^2 \cos^2 \check{\theta} + 2i\tilde{\epsilon} Fr \check{K} \cos \check{\theta}},
\end{aligned} \tag{7}$$

127 where the Jacobian of the coordinate transformation in Eq. (6) is $\frac{\check{K}}{W}$, and $\check{P}(\check{K}) =$

128 $\frac{\hat{P}(K_X, K_Y)}{W}$ to account for the reduced dimension. Simplifying,

$$\begin{aligned}
Z(\check{R}, \check{\varphi}) = & \\
129 \quad & - \lim_{\check{\epsilon} \rightarrow 0} \int_{-\pi/2}^{\pi/2} \int_0^{\infty} \frac{\sqrt{W^2 \cos^2 \check{\theta} + \sin^2 \check{\theta}} \check{P}(\check{K}) e^{-i\check{K}\check{R} \cos(\check{\theta} - \check{\varphi})} \check{K} d\check{K} d\check{\theta}}{\sqrt{W^2 \cos^2 \check{\theta} + \sin^2 \check{\theta}} - WFr^2 \check{K} \cos^2 \check{\theta} + 2i\check{\epsilon}FrW \cos \check{\theta}}. \quad (8)
\end{aligned}$$

130 The Sokhotski-Plemelj formula (Appel 2007) states that for a function $f(x)$ with
131 Cauchy principal value Q ,

$$132 \quad \lim_{\check{\epsilon} \rightarrow 0} \int_a^b \frac{f(x)}{x \pm i\check{\epsilon}} dx = \mp i\pi f(0) + Q \int_a^b \frac{f(x)}{x} dx. \quad (9)$$

133 Eq. (8) is manipulated into this form using

$$\begin{aligned}
x = \frac{\sqrt{W^2 \cos^2 \check{\theta} + \sin^2 \check{\theta}} - WFr^2 \check{K} \cos^2 \check{\theta}}{2FrW \cos \check{\theta}} \quad dx = -\frac{Fr \cos \check{\theta}}{2} d\check{K} \\
134 \quad f(x) = \check{K} \frac{\sqrt{W^2 \cos^2 \check{\theta} + \sin^2 \check{\theta}} \check{P}(\check{K}) e^{-i\check{K}\check{R} \cos(\check{\theta} - \check{\varphi})}}{Fr^2 W \cos^2 \check{\theta}}. \quad (10)
\end{aligned}$$

135 The zero point in f is determined uniquely by \check{K} since it is the only independent
136 variable. Therefore, designate

$$137 \quad \check{K}_0 \equiv \frac{\sqrt{W^2 \cos^2 \check{\theta} + \sin^2 \check{\theta}}}{WFr^2 \cos^2 \check{\theta}} \quad (11)$$

138 by solving with $x = 0$ in \check{K} . It is argued in Benzaquen et al. (2014) that
139 $Q \int_a^b f(x) dx/x$ is a rapidly decreasing function that can be ignored. Therefore,
140 the sea surface height is given by

$$141 \quad Z(\check{R}, \check{\varphi}) \approx i\pi \int_{-\pi/2}^{\pi/2} \frac{\check{P}(\check{K}_0) \left(W^2 \cos^2 \check{\theta} + \sin^2 \check{\theta} \right) e^{-iK_0 R \cos(\check{\theta} - \check{\varphi})}}{W^2 Fr^4 \cos^4 \check{\theta}} d\check{\theta}. \quad (12)$$

142 The pressure distribution is assumed to be of the form

143
$$p(x, y) = \frac{\pi mg}{Wb^2} e^{-\left(\frac{x\pi}{b}\right)^2 - \left(\frac{y\pi}{Wb}\right)^2}. \quad (13)$$

144 The leading coefficient is chosen such that the integral of the pressure function
 145 Eq. (13) over the plane of the sea surface is equal to the total force of the vessel
 146 on the water, mg (i.e., $p(x, y)$ has a net 'weight' equal to that of the vessel.)

147 The b and W parameters can be tuned to approximate the geometry of the
 148 physical hulls. To that end, $p(x, y)$ may be converted into an equivalent $\zeta(x, y)$
 149 to represent the sea surface elevation at which that pressure is experienced:

150
$$\zeta(x, y) = \frac{p(x, y)}{\rho g}. \quad (14)$$

151 The volume between $\zeta(x, y)$ and the xy axis defines a 3-dimensional figure, S that
 152 can be geometrically compared to a CAD hull model, C . In subsequent analysis,
 153 the W and b parameters are chosen such that the absolute norm $\int_{\mathbb{R}^3} |S - C| dV$ is
 154 minimized. The pressure distribution then matches the weight of the vessel and
 155 the shape is tuned to match that of the physical hull, within the constraints of
 156 the analytical model.

157 The physical pressure distribution is then converted into a form compatible
 158 with Eq. (12). In the nondimensional form specified by Eq. (4),

159
$$P(X, Y) = \frac{\pi m}{W\rho b^3} e^{-(X\pi)^2 - \left(\frac{Y\pi}{W}\right)^2}. \quad (15)$$

160 After a Fourier transform,

161
$$\hat{P}(K_X, K_Y) = \frac{m}{2\pi\rho b^3} e^{-\frac{1}{4\pi^2}(K_X^2 + W^2 K_Y^2)}. \quad (16)$$

162 And finally, applying the elliptic transform Eq. (6),

$$163 \check{P}(\check{K}) = \frac{m}{2\pi\rho b^3} e^{-\frac{\check{K}^2}{4\pi^2}}. \quad (17)$$

164 **Measurements**

165 On 17 September 2014 and 18 November 2014, the wakes of the ferry boats
166 servicing the Seattle-Bremerton route were measured via data buoys deployed
167 from R/V Jack Robertson. An approximate sailing line for an approaching ferry
168 was obtained from Automatic Identification System (AIS) vessel tracking data.
169 Buoys were then deployed in a line running approximately perpendicular to the
170 predicted sailing line of the ferry. The ferry crossing time was recorded as the
171 time when the AIS coordinates of the ferry intersected the time-dependent best
172 fit line of the array of buoys.

173 The data buoys in use consisted of four SWIFT and eight μ SWIFT buoys
174 (Thomson 2012). These buoys are free-floating spar-type buoys equipped with
175 GPS loggers. The μ SWIFT uses a QStarz BT-Q1000eX and the SWIFT uses a
176 Microstrain 3dm-35. Data were sampled at rates varying from 4 Hz to 10 Hz.
177 The GPS logs both positional velocity data, but the horizontal positional ac-
178 curacy of commercial GPS devices such as this one are only on the order of 10m–
179 significantly larger than is required to phase-resolve individual waves. The rel-
180 ative horizontal velocity resolution, however, is approximately 0.05 m/s through
181 Doppler phase processing of the raw signals (Herbers et al. 2012; Thomson 2012).
182 The vertical elevation and velocity do not have sufficient accuracy to be consid-
183 ered for data processing. Therefore, the positional GPS data are used to obtain
184 approximate locations and trajectories of the buoys, and the horizontal velocity
185 data are used to capture the motion of the sea surface. Typically, buoy motion
186 due to ocean current was less than 1 m/s.

187 In linear wave theory, the motion of inertial particles influenced by sinusoidal
188 wave motion is given by

$$189 \quad v(t) = a\omega e^{kz} \cos \theta(t) \quad (18)$$

190 for wave amplitude a , angular frequency ω , wavenumber k , and distance below
191 the surface z . Similarly, the surface elevation is given by

$$192 \quad \zeta(t) = a \cos \theta(t). \quad (19)$$

193 For a particle on the surface, the sea surface height may be obtained from the
194 velocity via the relation

$$195 \quad \zeta(t) = \frac{v(t)}{\omega(t)}. \quad (20)$$

196 The natural frequency of both types of drifter buoy is well above the frequencies
197 of the observed wake pattern. As such, the drifters are treated as inertial surface
198 particles. The goal of the time domain analysis is, therefore, to obtain accurate
199 and well-behaved values for ω and $v(t)$ in order to reconstruct the sea surface
200 height. Here, well-behaved refers to avoiding low-frequency velocity drift and
201 artificially small values of ω .

202 The surface reconstruction is performed by a process of

- 203 • Calculating and subtracting low-pass filtered velocity vectors, in order to
204 remove the effects of buoy drift
- 205 • Obtaining a scalar velocity by calculating the principal axes for buoy kinetic
206 energy in the plane of the ocean and taking only the velocity of the major
207 axis
- 208 • Calculating a local wave frequency ω by calculating the zeros in the scalar
209 velocity
- 210 • Applying Eq. (20) to obtain the sea surface height. The conversion from

211 scalar velocity to sea surface height is shown in Fig. ??

212 In a rough sense, the surface disturbance caused by a passing ferry boat can
 213 be expressed as a near-Dirac disturbance, or equivalently, a broadband wave gen-
 214 eration. Denote such a function

$$215 \quad \zeta(x, 0) = \int_{-\infty}^{\infty} a(k) e^{ikx} dk \quad (21)$$

216 where $a(k)$ specifies a magnitude associated with a given wavenumber k . Then,
 217 for the outward spreading of this disturbance, the dispersion relation $\omega(k) = \sqrt{gk}$
 218 is used. This is the deep water limit, which for the shallowest in-situ depth
 219 encountered of 20 meters, is permissible for wavelengths of less than 40 meters.
 220 The time-dependent sea surface height is then

$$221 \quad \zeta(x, t) = \int_{-\infty}^{\infty} a(k) e^{i(kx - \sqrt{gk}t)} dk. \quad (22)$$

222 In the limit of large x and t (as in the case of the field data), a stationary
 223 phase approximation (Stoker 1957) may be invoked in order to evaluate Eq. (22),
 224 yielding

$$225 \quad \zeta(x, t) \approx a(k_0) \sqrt{\frac{gt^2}{\pi x^3}} \cos \left(k_0(x) - \omega(k_0)t \pm \frac{\pi}{4} \right) \quad (23)$$

226 where $k_0 = gt^2/(4x^2)$ solves the equation $\partial/\partial k [kx - \sqrt{gk}t] = 0$.

227 Therefore, for a given (x, t) displacement from the initial disturbance (i.e., a
 228 time t after a vessel's closest approach to a point x of observation), the observed
 229 frequency of waves should be given by

$$230 \quad \omega(x, t) = \frac{gt}{2x}. \quad (24)$$

231 Note that in Eq. (24) the expression for velocity is the group velocity, $g/(2\omega)$,

232 rather than the phase velocity. For constant $x = d$ (i.e. a buoy that is stationary
 233 in the reference frame of the ocean surface) the observed frequency should in-
 234 crease linearly in time with slope $g/(2d)$. Comparisons of predicted and observed
 235 wave frequencies (Fig. ?? being a typical case) show that treating the vessels as
 236 generating broadband wave pulses serves as a good predictor for the arrival of
 237 wave frequencies observed by the buoys. This is a useful check in cases where the
 238 distance from sailing line is difficult to determine.

239 RESULTS

240 Analytical Model

241 The analytical model described by Eq. (12) and Eq. (13) produces a steady
 242 state sea surface height as a function of boat length b , boat aspect ratio W , boat
 243 mass m , water density ρ , and the boat velocity. If Eq. (12) is expressed with the
 244 leading coefficients in dimensional form, the expression becomes

$$245 \quad \zeta(x, y) = \frac{m}{4\pi^2 \rho b^2} \text{Im} \left[\int_{-\pi/2}^{\pi/2} \frac{e^{-\frac{K_0}{4\pi^2}} \left(W^2 \cos^2 \check{\theta} + \sin^2 \check{\theta} \right) e^{-iK_0 R \cos(\check{\theta} - \check{\varphi})}}{W^2 \text{Fr}^4 \cos^4 \check{\theta}} d\check{\theta} \right] \quad (25)$$

246 One may conclude, then, that the sea surface height scales linearly as the boat
 247 mass, and inversely as the water density. Eq. (25) contains a complex dependence
 248 on the remaining variables. These remaining dependencies are shown graphically
 249 in Fig. ?. The sea surface forms a distinctive Kelvin wake pattern. In Fig. ?,
 250 two such solutions are shown, for different values of W . At low W , the largest
 251 waves are located at the intersection between the angled diverging wake and the
 252 approximately horizontal transverse wake. At $W = 1$, indicating a circular pres-
 253 sure distribution, the greatest perturbation in the sea surface is located directly
 254 behind the vessel, in the centers of the transverse wake waves. The growth of the
 255 diverging wake relative to the transverse wake is typically associated with a larger

256 vessel Froude number (higher velocity, or smaller vessel length when all else is
257 held constant). The effects of this transition are further explored on page 16.

258 The plots in Fig. ?? explore the maximum wake height, as this is the relevant
259 parameter under study for the ferry boats. Maximum wake height is defined to
260 be the absolute value of the greatest deformation in the free surface, at a given
261 perpendicular distance d from the axis of vessel motion. Of the variables other
262 than m and ρ , the maximum wake height is determined by the Froude number and
263 the aspect ratio W . However, because the Froude number contains two physically
264 relevant variables, velocity and characteristic length, their independent effects
265 are explored. In each case, in Fig. ??, the vessel mass and water density are
266 unchanged.

267 In order to address the physical wakes, the model parameters were chosen to
268 represent an Issaquah class and a Super class vessel. Using the available vessel
269 dimensions, mass, and CAD models, Gaussian pressure distributions of the form in
270 Eq. (13) were created. The resulting wake profiles, shown in Fig. ?? and generated
271 from the values in Table ??, were highly dependent on the boat length b , but in all
272 cases, the Issaquah class model produced much larger wakes than the Super class
273 model. The presented models match velocity rather than Froude number, for dual
274 reasons. The ferry operating conditions are determined by the time schedule, and
275 so the same velocity must be attained when comparing performance. In addition,
276 the Froude number of the Issaquah class is larger by a small enough amount that
277 slowing the vessel to attain Froude number parity would still not eliminate the
278 difference in wake height. The wake profiles, presented at the same scale, indicate
279 significantly smaller wake for the Super Class, as well as a weaker diverging wake
280 relative to the transverse wake.

281 **Measurements**

282 A total of 19 ferry crossings were recorded, yielding 79 useful wake histories
283 from the drifter buoy data. These wake histories represent 65 measurements of
284 the Issaquah class at a variety of vessel speeds, and 14 measurements of the Super
285 class, all at approximately its standard operating speed of 18 knots.

286 For one wake event, a μ SWIFT and a SWIFT buoy were placed within 10m of
287 each other in order to compare buoy performance. In general, the SWIFT buoy
288 provides less noisy data, but the comparison in Fig. ?? suggests that both varieties
289 of buoy have similar spectra at the frequencies of interest.

290 The raw buoy velocities also indicate close agreement. However, the sea surface
291 height reconstruction is highly sensitive to the exact recorded velocities. Despite
292 qualitatively similar buoy velocities, the final sea surface height differs consider-
293 ably between the two measurements. This reflects the highly localized nature of
294 sea surface velocities. The spread in maximum sea surface height at consistent
295 experimental conditions (see Fig. ??) is qualitatively similar for both buoy types
296 when the aggregate data are taken. For this reason, data from both buoy types
297 are used interchangeably.

298 The angular frequency associated with the maximum wake amplitude for a
299 wake event was recorded. In general, the observed frequency was lower at larger
300 distances from the sailing line, as the dispersion relation for deep water waves
301 would suggest. However, the relationship between angular frequency and max-
302 imum wake height, Fig. ??, demonstrated a trend that appears to depend on
303 the vessel speed. The most energetic waves are the ones with low angular fre-
304 quency and large maximum height. Therefore, it is particularly concerning that
305 the largest wakes associated with higher vessel speeds are also the lowest frequency
306 wakes.

307 **Comparisons**

308 Because one of the primary concerns for WSDOT ferry wakes is the overtop-
309 ping of sea walls, and the shoreline morphology is quite steep, the maximum sea
310 surface height is considered as a metric for comparison. Maximum sea surface
311 height versus distance from sailing line was compared across all measurements
312 and models. The analytical model’s parameters were tuned to match the shape
313 and mass of the Issaquah and Super Class hulls, as described on page 4. Fig. ??
314 displays the overall comparison between the various wake measurements studied.
315 The analytical model does a good job of capturing the decay in wake height with
316 increased distance from the sailing line that is observed in the field data. At low
317 vessel speed, however, the analytical model tends to predict a smaller disturbance
318 than is physically observed.

319 Another method of visualizing the model’s predictive ability is to treat each
320 buoy as an inertial particle moving on the surface of the ocean, in the reference
321 frame of the vessel. The buoy’s location is converted to a time varying displace-
322 ment from the vessel, and the analytic solution is calculated along the transect
323 the buoy takes, parameterized by time. The parameterized analytical solution
324 $\zeta(t)$ then mimics the history of the sea surface height experienced by a numerical
325 buoy taking the same trajectory as the buoy in the field. This permits direct
326 comparisons with reconstructed sea surfaces from the wave buoys.

327 Several such comparisons are presented in Fig. ?.?. This method is quite sensi-
328 tive to clock synchronization between the vessel and buoy, as well as vessel speed.
329 For each comparison, the vessel class and speed are matched between the ana-
330 lytical model and the observations for the measured vessel crossing. The initial
331 wake structure is well-captured, though the timing may have an offset due to
332 ocean currents and the exact location of the AIS receiver on the ferry. The wave
333 packeting that is often seen by the physical buoys was not reliably reproduced by

334 the model, and the reasons for this are explored on 16.

335 DISCUSSION

336 Models and data indicate that at similar operating speeds, the Issaquah Class
337 vessels make significantly larger wakes than the Super Class vessels. This result is
338 somewhat unintuitive, as the Super Class is a longer and heavier vessel. Results
339 from the analytical model point to a dual justification for the wake height disparity.
340 The increased vessel length for the Super Class has the effect of reducing the vessel
341 Froude number, and the decreased aspect ratio W also works to decrease the wake.
342 These two effects counteract the increased weight to yield an overall lower wake.

343 Furthermore, wake height reduction is achieved by reducing the vessel speed
344 for the Issaquah Class. The Super Class data do not include significant speed
345 variation, but the data displayed in Fig. ?? demonstrate that the full-speed Su-
346 perclass produces much smaller peak wake height than the full-speed Issaquah,
347 and qualitatively comparable wake heights to the speed-limited Issaquah class.
348 The field data support the conclusions obtained in the 1985 study (Nece et al.
349 1985). Analytic data indicate a roughly linear relationship between peak wake
350 height at a given distance from the sailing line and vessel speed in the regime of
351 interest. Therefore, in assessing a given distance and peak height objective, the
352 field data plots should be used as two data points, with a linear interpolation to
353 determine appropriate speed.

354 At low Froude numbers, the transverse wake dominates the wake field of a ves-
355 sel. However, field observations and CFD modeling have shown that, particularly
356 at low Froude numbers, the vessels produce a significant bow wake. The bow wake
357 is a consequence of the physical displacement of the water in front of the vessel
358 and is not captured by the model. Particularly at the lowest operational speeds
359 of the ferries, the bow wake dominates the waves produced by the vessel. This is
360 the proposed mechanism for the model failure at low Froude number (below 0.35).

361 In addition, the packeting observed in Fig. ?? may be a beat frequency produced
362 by the interaction of the bow wake and the Kelvin wake. At $Fr > 0.35$, the bow
363 wake was not observed to affect the maximum wake height measurements.

364 CONCLUSIONS

365 The wakes produced by Issaquah and Super Class WSF vessels were studied
366 using an array of GPS buoys and with an analytic potential flow wake model.
367 Each of these methods produced results that indicate a maximum wake height
368 for a given distance from the sailing line of the vessel and a given vessel speed.
369 These results are summarized in Fig. ?.?. The in-situ buoy measurements are
370 regarded to be the most reliable, but also the contain the most sparse data, as
371 well as significant noise. A higher-fidelity sea surface height reconstruction may
372 be possible using vertical acceleration, given a high enough sampling rate. The
373 4-10 Hz sampling rate used presently was insufficient to accurately represent buoy
374 heave.

375 Despite being an older and larger vessel, the Super Class ferry produces smaller
376 wakes than the Issaquah class at a given speed. The relatively simple analytic
377 representation of these vessels is able to attribute difference in wake height to the
378 geometric qualities of aspect ratio and vessel length, as well as vessel mass. The
379 measurements and subsequent model comparison constitute a novel approach the
380 question of vessel wake management.

NOTATION

Variable Name	Description
ρ	Seawater density. Assigned a value of $1025kg/m^3$ for Puget Sound.
ω	Angular frequency of a wave.
$\{x, y, z\}$	Cartesian coordinates in the frame of the ferry boat. x denotes the streamwise direction, y is the crossstream direction, and z is the vertical direction. The origin of the grid is located at the mass centroid in x and y , and at the flat sea surface in z .
k_x, k_y	Wavenumbers associated with the analytic wake model.
W	Aspect ratio of a ferry boat in the analytic wake model. This is the ratio of the major to minor axis of the ellipse representing the boat. A smaller W implies a more streamlined vessel.
b	Characteristic length of the ferry.
Fr	Length Froude number, defined by $U/(\sqrt{gb})$.
U	Free stream velocity of flow past the ferry boat.
ζ	Value of the sea surface elevation, measured as a signed perturbation from a flat surface of value 0.
Z	Nondimensional sea surface, defined by $4\pi^2\zeta/b$.
g	Gravitational acceleration, given a value of $9.8m/s^2$.
h	Maximum wake height. This is the maximum of the absolute value of the perturbation in the free surface.
a	Wave amplitude, for the purpose of linear wave analysis.

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